

# Macroscopic lithotype characterisation of the 1st Middle-Polish (1st Lusatian) Lignite Seam in the Miocene of central Poland

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## Abstract

The 1st Middle-Polish (1st Lusatian) Lignite Seam is exploited in open-cast mines in central Poland. A large number of lignite lithotypes, grouped in four lithotype associations, are distinguished: xylic, detritic, xylo-detritic and detro-xylic lithotype associations, which show various structures. Each lithotype association was produced under specific peat-forming environmental conditions. In the case of the lignite seams under study they represent all the main environments that are known from Neogene mires, i.e.: fen or open water, bush moor, wet forest swamp and dry forest swamp.

For a simple and practical description in the field of both the lignite sections and borehole cores, a new codification for lignite lithotypes is proposed. It is based on the codification of clastic deposits (lithofacies). The practical value of the new lignite lithotype codification is examined in three vertical sections of the 1st Middle-Polish Lignite Seam.

**Keywords:** lignite lithotypes, sedimentology, First Middle-Polish Lignite Seam, Miocene, central Poland

## 1. Introduction

The 1st Middle-Polish (1st Lusatian) Lignite Seam (MPLS-1) is exploited only in the surroundings of Konin and Turek. The Lubstów open-cast mine (Fig. 1), where the 2<sup>nd</sup> Lusatian Lignite Seam has been extracted, was closed in 2009. The annual lignite production in this region is about 13–14 million tones, and almost 100% is consumed by electrical power plants.

During the last 70 years, over 7000 boreholes have been drilled to document the lignite deposits in the study area (Widera, 2007; Widera et al., 2007). Moreover, every year additional boreholes are drilled for detailed petrographical analyses, and the walls of the excavations are also mapped. These activities require de-

scriptions in a proper form, but this is complex and time-consuming without a proper lithotype codification. Previously used codes for lignite lithotypes were limited to textural features only (cf. Brzyski, 1981; Kwiecińska & Wagner, 1997; Taylor et al., 1998; Markič et al., 2001).

### 1.1. Objectives

The main objective of the present contribution is to distinguish the lithotype compositions of the MPLS-1, which has been exploited because of its extensive presence in central Poland. The distinguishing between lithotypes should show the relationship between peat-

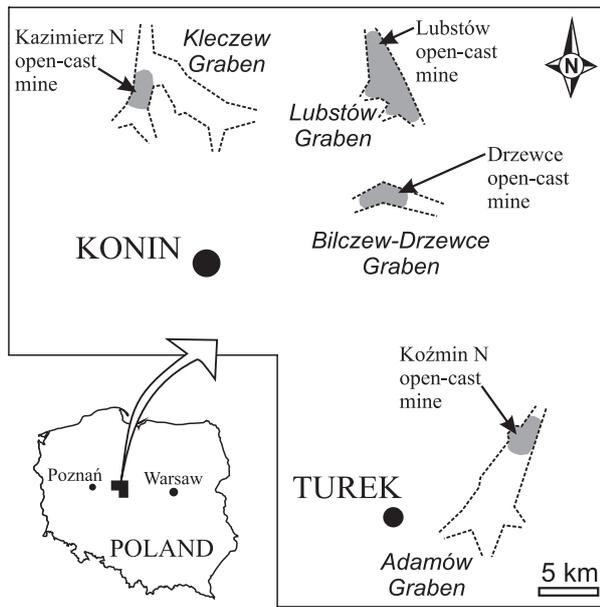


Fig. 1. Location map of the lignite open-cast mines in the grabens of the Konin-Turek area, central Poland.

forming environments and the resulting lignite lithotypes.

A new codification for lignite lithotypes has been developed for the purpose. It is based on the combination of textural and structural properties of the lignite seam. The proposed lithotype codes should be applicable in the field and during the documentation of sections of the lignite seam.

## 1.2. Geological setting

In central Poland, the major lignite deposits are located in grabens that were tectonically active in the Paleogene and/or Neogene (Kasiński & Piwocki, 1992; Widera, 2007; Widera & Hałuszczak, 2011). The lignites under study here are exposed in the Drzewce, Kazimierz N and Koźmin N open-cast mines, which occur in the Drzewce-Bilczew, Kleczew and Adamów grabens, respectively (Fig. 1). The Drzewce-Bilczew and Adamów grabens started their development at the Eocene/Oligocene boundary (Widera & Kita, 2007; Widera, 2010). Since the Early Oligocene, the entire area of central Poland was subjected to uplift (Widera et al., 2008; Jarosiński et al., 2009; Widera & Hałuszczak, 2011). In contrast, the Kle-

czew Graben started to subside at the turn of the Oligocene/Miocene, as shown by the presence of Miocene sediments at the base of the succession (Widera, 2007).

In all cases the MPLS-1 is underlain by sands with intercalations of silts and lignites of Early-Middle Miocene age. The sands are thick: from a few decimetres to up 50 m. The lignites of the MPLS-1 show an average thickness of 6.5–8.7 m (Sadowska & Giża, 1991; Ważyńska, 1998). Their age is middle Middle Miocene (Piwocki & Ziemińska-Tworzydło, 1997), so that they correlate in time with the 1<sup>st</sup> Lusatian Lignite Seam of SE Germany (Standke et al., 1993; Grimm, 2002). The Polish lignite is a low-rank coal (reflectance,  $R_r = 0.26\text{--}0.29\%$ ; Kwiecińska & Wagner, 1997, 2001) with an ash content of 11.0–12.0 wt% and a sulphur content of 0.8–0.9 wt%. These properties of the lignites from the Konin-Turek region make them suitable as a cheap fuel for the production of electricity for decades (Piwocki, 1987; Kasiński et al., 2006; Bielowicz, 2010).

The MPLS-1 is overlain by fine-grained deposits, mainly clays and silts with lenses of fine sands. Their total thickness ranges from a few metres to more than 30 m. They were probably deposited in an anastomosing-river environment (Widera, 2007) from the middle part of the Middle Miocene to the lowermost part of the Early Pliocene (Piwocki & Ziemińska-Tworzydło, 1997).

## 2. Former nomenclature and codification of lignite lithotypes and lithofacies

The currently used lignite lithotype terminology and classification is highly complex (cf. Brzyski, 1981; Markič et al., 2001) and differs from the ICCP (1993) recommendations (Taylor et al., 1998). Until now the term 'lithotype' has been used in textural descriptions of lignites. In a sedimentological sense, however, a lignite lithotype is a macroscopically recognisable and homogeneous stratum of low-rank coal with a thickness of more than 5 cm (Kwiecińska & Wagner, 1997; Ticleanu et al.,

**Table 1.** Codes for the main lignite lithotypes according to Kolcon & Sachsenhofer (1999).

lithotype	code
xylic coal	X
detrilic coal	D
detro-xylic coal	D-X
xylo-detrilic coal	X-D
fusilic coal	F

**Table 2.** Codes for the main lignite lithotypes according to Ticleanu et al. (1997).

lithotype	code
coal	XC
detrilic coal	DC
weak coal	WXC
clayey coal	CC
charcoal	CH

1999). Two lithotype codifications, which are still the most popular, are presented in Tables 1 and 2. These codes are applicable to lignite seams only. Much confusion is therefore raised when sections containing both lignite seams and associated clastic deposits are described.

The lithofacies code for the description of sediments that was established by Miall (1977) and then slightly modified (e.g. Eyles et al., 1983; Ghibaudo, 1992; Rhee & Chough, 1993; Mastalerz, 1995; Zieliński, 1995; Krüger & Kjær, 1999), is widely used. This code consists of letter symbols for grain size (e.g. G denotes gravel, S = sand, F = fines = silt + clay, D = diamict, C = organic deposits, etc.) and structure (m = massive, h = horizontally stratified, p = planar cross-stratified, d = deformed, etc.). For example, the lithofacies Gh and GSh can be denoted as horizontally stratified gravel and horizontally stratified sandy gravel, respectively.

A problem arises, however, when the lithotype nomenclature and lithofacies codes are used together. The 'F' code can refer to 'fusilic coal' in the lignite lithotype codification (Table 1; Kolcon & Sachsenhofer, 1999) or can be identified as 'fines', i.e. silt and clay, in the lithofacies codification (Miall, 1977). Similarly, the 'D' code can mean 'detrilic coal' in the lignite lithotype codification (Table 1; Markič & Sachsenhofer, 1997) or can be identified as 'diamict'

in the lithofacies codification (Eyles et al., 1983). To avoid any possible confusion between these codes, a new codification of lignite lithotypes is therefore proposed.

### 3. Proposed nomenclature and codification of lignite lithotypes

A change of the definition of the term 'lithotype' is proposed. A term that is well applicable in practice should describe – as done in the lithofacies nomenclature (Miall, 1977) – both textural and structural features of lignite. The term 'lignite lithotype' should therefore refer to a macroscopically recognisable and homogeneous stratum of low-rank coal (lignite) with a thickness of more than 5 cm, which is characterised by a specific texture and structure.

The most adequate for the lignite texture is the nomenclature established by Wolf (1988), and then adapted and widely propagated by Markič & Sachsenhofer (1997) and Kolcon & Sachsenhofer (1998, 1999). The codes for the lignite texture, used by these authors, should, however, be changed for the above reasons (Table 1). It is proposed here to use the capital letters: X (for xylic), D (for detrilic), XD (for xylo-detrilic), DX (for detro-xylic), and F (for fusilic), always in combination with a capital letter L (lignite). For example, the 'XDL' code refers to a 'xylo-detrilic lignite' texture in this lignite lithotype codification (Table 3).

Moreover, the internal structure of lignite should be coded with lower-case letters, again following the sedimentary lithofacies codification. Field experience shows that four letters are enough for the basic description of the lignite structure. Appropriate lower-case letters for the structure of lignites are: m (for massive), h (for horizontally stratified), g (for gelified), and d (for deformed). Deformed lignite can be coded by two extra letters (between brackets) to indicate the type of deformation: (fr) for fractured, (fa) for faulted, and (fo) for folded. For example, code 'XLm' refers to the 'xylic lignite with a massive structure' code 'DLd(fo)' indicates a 'detrilic lignite with a deformed (folded) structure' (Table 3).

**Table 3.** Proposition for a new lignite-lithotype codification (nomenclature partly by Wolf, 1988).

lithotype texture	code
xylitic lignite	XL
detritic lignite	DL
detro-xylitic lignite	DXL
xylo-detritic lignite	XDL
fusitic lignite	FL

lithotype structure	code
massive	m
horizontally stratified	h
deformed (fractured, folded, faulted)	d (fr, fo, fa)
gelified	g

#### 4. General description of the lignite lithotype associations

The classification of the lignite lithotypes is based on a quantitative criterion, the so-called '10% rule'. It is used by most scientists and also recommended by the ICCP (1993). The content of such a lignite-seam elements is taken into consideration as xylitic fragments, fine-detrital matrix and/or fusitic particles (e.g. Wolf, 1988; Kwiecińska & Wagner, 1997; Markič & Sachsenhofer, 1997; Taylor et al., 1998; Kolcon & Sachsenhofer, 1998, 1999). Obviously, each of described lithotypes may represent several possible combinations of structures in the codes listed below (cf. Table 3).

##### 4.1. Xylitic lignite (XL) association

This lignite lithotype association contains at least 90% of xylites, i.e. well fossilized wood fragments such as stumps, trunks and branches. They must be larger than 1 cm and, therefore, the interior structure of wood is macroscopically well visible. The matrix of this lithotype association consists of fine-detrital (< 1 cm) matter. In most cases the xylites lay horizontally (e.g. trunks, branches, twigs), but sometimes they are in growth position (e.g. stumps, roots). For this lithotype association, all structures (i.e. massive, horizontally stratified, gelified and deformed) are characteristic

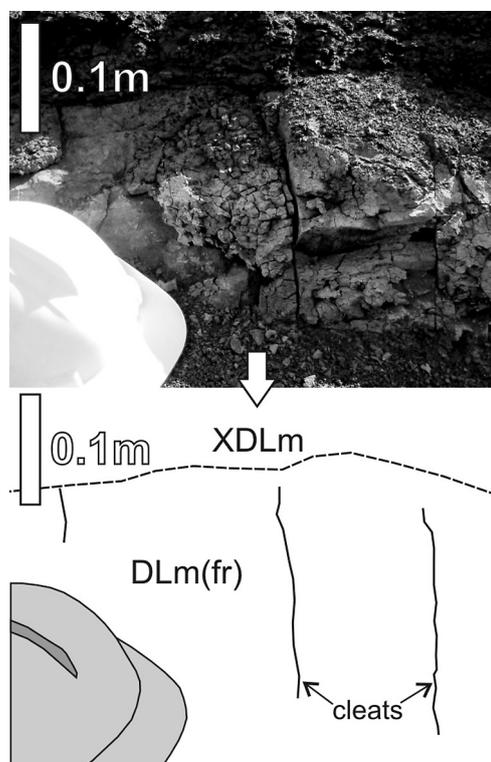
and included in the lignite lithotype codification presented here (Table 3).

##### 4.2. Detritic lignite (DL) association

This association consists of more than 90% of fine-graded (< 1 cm), usually homogeneous organic matter. The rest (up to 10%) of the components are xylitic fragments. Detritic lignite is mostly characterised by a massive and/or deformed structure, e.g. with cleat-like fractures (Fig. 2), but sometimes it can be horizontally or pseudo-horizontally stratified. A gelified structure is sporadically present in detritic lignite.

##### 4.3. Xylo-detritic lignite (XDL) and detro-xylitic lignite (DXL) associations

These two lithotype associations are a combination of both the xylitic and detritic ones. They have various structures. Paradoxically, it has been historically established that xylo-



**Fig. 2.** Fractures developed within the xylo-detritic lignite lithotype (Kazimierz N open-cast mine). For explanation of the lignite lithotype codes, see Table 3.

detritic lignite contains more detrital matrix than xylites. In contrast, the detro-xylitic lignite comprises more xylitic particles than the surrounding fine-detrital matrix (cf. Brzyski, 1981; Wolf, 1988; Kwiecińska & Wagner, 1997; Markič & Sachsenhofer, 1997; Taylor et al., 1998; Kolcon & Sachsenhofer, 1999; Markič et al., 2001). On the other hand, the sum of the xylites and detrital matrix should be more than 90% (the '10% rule' – in the described two lignite lithotypes). These lithotype associations

represent all structures included in Table 3, e.g. a folded structure (Fig. 3).

#### 4.4. Fusitic lignite with massive structure (FLm) association

The presence of the fusitic lignite association in a lignite seam indicates that fires happened in a mire (e.g. Kwiecińska & Wagner, 1997; Ticleanu et al., 1999). Fusitic lignite is actually a fossil charcoal; it is, therefore, relatively rare in a lignite-seam profile. Fusitic lignite is most commonly a part of other lithotypes. It forms lenses rather than continuous strata, or it occurs as encrustations around xylites. This lithotype association is black or silky-black, highly porous and friable, with needle-shaped particles (Wolf, 1988; Markič & Sachsenhofer, 1997; Kolcon & Sachsenhofer, 1999). The macroscopically visible structure can consequently only be massive.

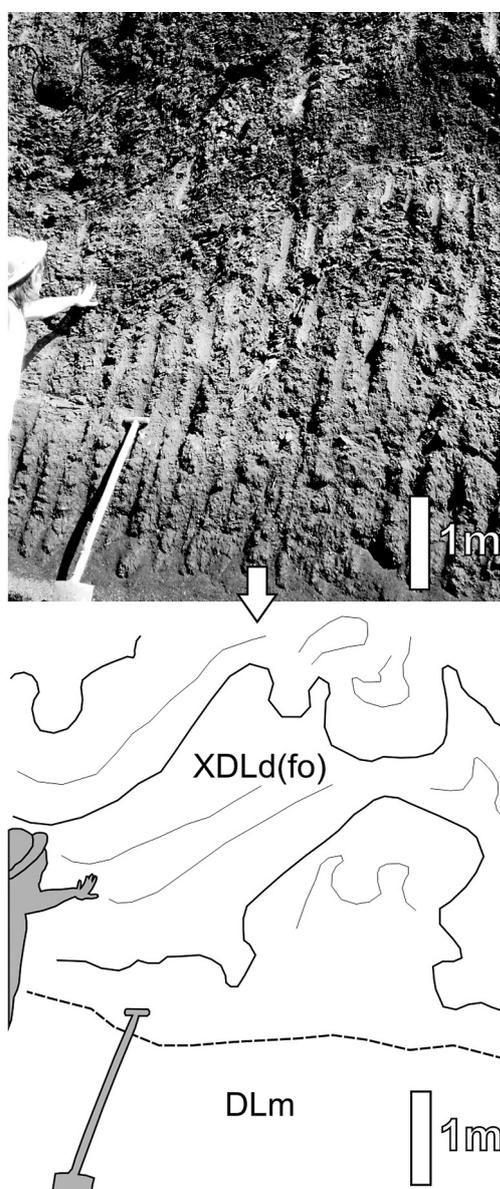
### 5. Macroscopic description of the lignite seam under study

The characteristics of the MPLS-1 can be studied in three vertical sections from the lignite open-cast in the surroundings of Konin and Turek, viz. Kazimierz N, Drzewce and Koźmin N (Fig. 1).

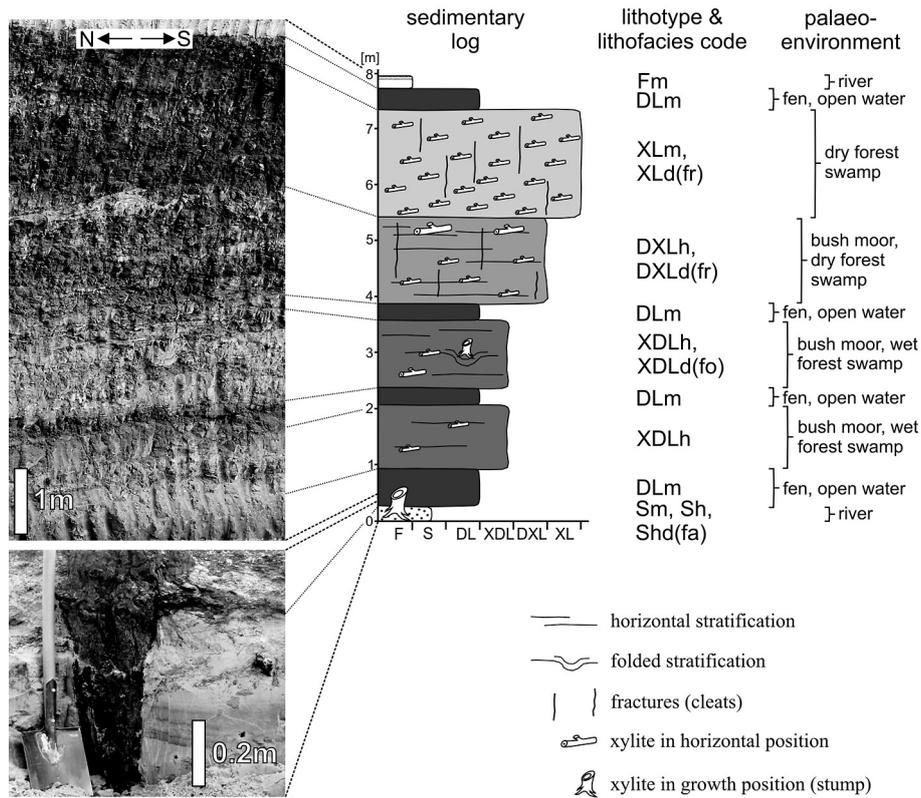
#### 5.1. The Kazimierz N site

The Kazimierz N section is exposed in the northern part of the Kazimierz N open-cast mine. It belongs to the western, S-N-trending branch of the Kleczew Graben, located ~20 km NNW of Konin (Fig. 1). The altitude of this section ranges between 44.2 and 52.2 m a.s.l., and the thickness of its lignite seam is 7.7 m only (Fig. 4).

Four lignite lithotypes associations and two lithofacies ones have been distinguished (Fig. 4). The latter comprise (1) lithofacies association Fm, Sm, and (2) lithofacies association Sh, Shd(fa). They represent the clastic deposits that, respectively, underlie and overlie the li-



**Fig. 3.** Folds in the xylo-detritic lignite lithotype (Drzewce open-cast mine). For explanation of the lignite lithotype codes, see Table 3.



**Fig. 4.** Macroscopic vertical section of the 1<sup>st</sup> Middle-Polish Lignite Seam in the Kazimierz N open-cast mine. Note the shallowing-upward trend characterised by a transition from detritic lignite at the base to xylitic lignite at the top. For explanation of the lignite lithotype codes, see Table 3; for explanation of the lithofacies codes, see the text.

nite seam. On the other hand, four lignite lithotype associations consist of one or two lignite lithotypes.

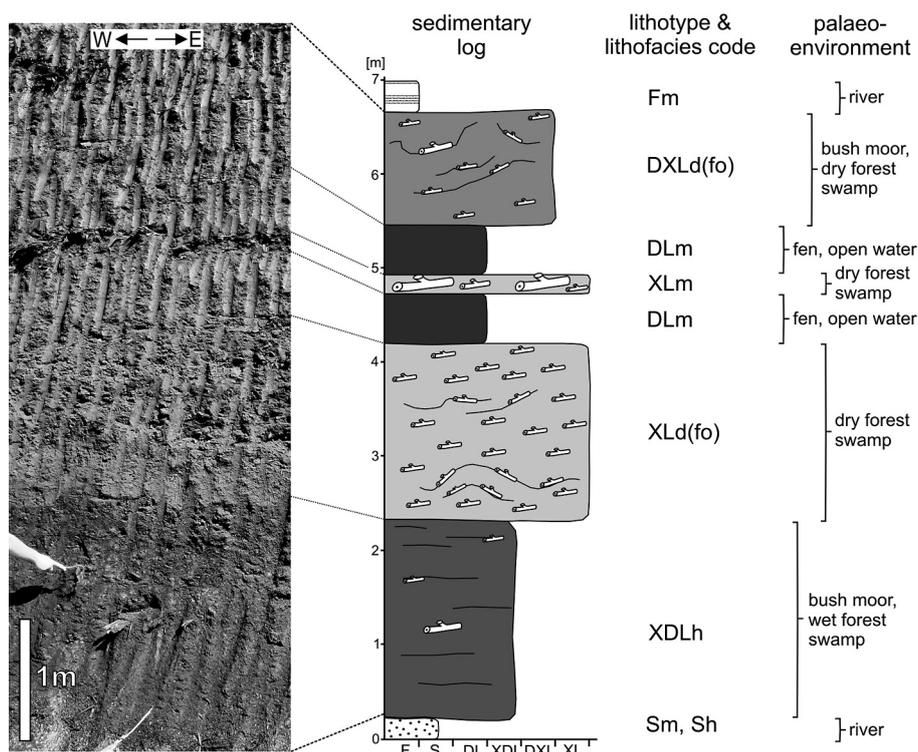
The section starts with detritic lignite with a massive structure, i.e. DLm, which is 0.7 m thick. Moreover, two layers of this lithotype are also situated between other lithotypes or lithotype associations: XDLh and XDLh, XDLd(fo) as well as XDLh, XDLd(fo) and DXLh, DXLd(fr) (cf. Table 3). The thickest (up to 1.8 m) lithotype association, consisting of lithotypes XLm or XLd(fr), is exposed in the upper part of the section. The section ends with detritic lignite with a massive structure, i.e. DLm, which is 0.3 m thick (Fig. 4).

Most of the lignite lithotypes are deformed (folded and/or fractured): XDLd(fo), DXLd(fr) and XLd(fr). Additionally, the presence of the xylites in growth position is also very characteristic in the vertical section of the lignite seam (Fig. 4).

## 5.2. The Drzewce site

This lignite section is situated in the central-northern segment of the Drzewce open-cast mine. It occurs in the axial part of the W-E trending Bilczew-Drzewce Graben, which is located ~15 km ENE of Konin (Fig. 1). The altitude of the section ranges from 61.8 m to 68.8 m a.s.l. The lignite seam is 6.6 m thick (Fig. 5).

A clastic lithofacies association underlies the lignite seam and is composed of massive as well as horizontally stratified sand, i.e. Sm and Sh. The lignite seam contains five lithotypes, which can be grouped in four lignite lithotype associations (Fig. 5). The horizontally stratified xylo-detritic lignite lithotype, XDLh, is the lowermost in the section. It is overlain by a 2.1 m thick lithotype of xylitic lignite with a slightly folded structure XLd(fo). Two layers of massive detritic lignite, DLm, follow, with intercalation of a xylitic horizon, XLm, of 0.2-0.3 thick. A detro-xylitic lignite with folded structure, DXLd(fo), forms the top part of the



**Fig. 5.** Vertical section of the 1<sup>st</sup> Middle-Polish Lignite Seam in the Drzewce open-cast mine. Note the presence of lithotypes with a large xylites content in the middle of the section. For explanation of the lignite lithotype codes, see Table 3; for explanation of the lithofacies codes, see the text; for graphic explanations, see Figure 4.

lignite seam. Massive fines (silt + clay), Fm, overlie the lignite seam in the section (Fig. 5).

### 5.3. The Koźmin N site

This section is located in the northern part of the Koźmin N open-cast mine, which is situated in the Adamów Graben. This SSW-NNE trending tectonic structure is located 5–20 km E and NE of Turek (Fig. 1). The altitude of the presented section ranges between 57.3 and 64.0 m a.s.l. The lignite seam is 6.2 m thick (Fig. 6).

Massive sand and/or fines (silt + clay) with a much organic matter, SCm and FCm, lay at the base of the section; it is noteworthy that xylites in growth position are very common in the clastic deposits under the lignite seam (Fig. 6). Then follows the lignite-bearing part, in which only three lignite lithotype associations are present (Fig. 6). The xylo-detritic lignite lithotype association, XDL, dominates; it shows various structures: XDLm, XDLh, XDLd(fr) and XDLd(fo). The layers belonging to this association jointly are 5.0 m thick. Ad-

ditionally, massive detritic lignite, DLm, and fractured detro-xilitic lignite, DXLd(fr), are present in the middle part of the lignite seam (Fig. 6). Massive fines, Fm, rest at the top of the section.

## 6. Interpretation and discussion

The MPLS-1 in central Poland is underlain and overlain by clastic sediments (Figs 4–6). Taking into consideration the depositional architecture, palaeogeographical position and lithofacies assemblages, these clastic deposits are characteristic of a river environment. The sediments under the lignite represent an alluvial fan and braided river, whereas those on top of the seam are characteristic of an anastomosing river (Widera, 2007, 2010).

The reconstruction of a mire type and resulting lignite lithotypes (Fig. 7) has been based on the pioneering studies by Teichmüller (1958). The relationship between the peat-forming environment and lignite lithotypes has been discussed by Wolf (1988), Teichmüller (1989),

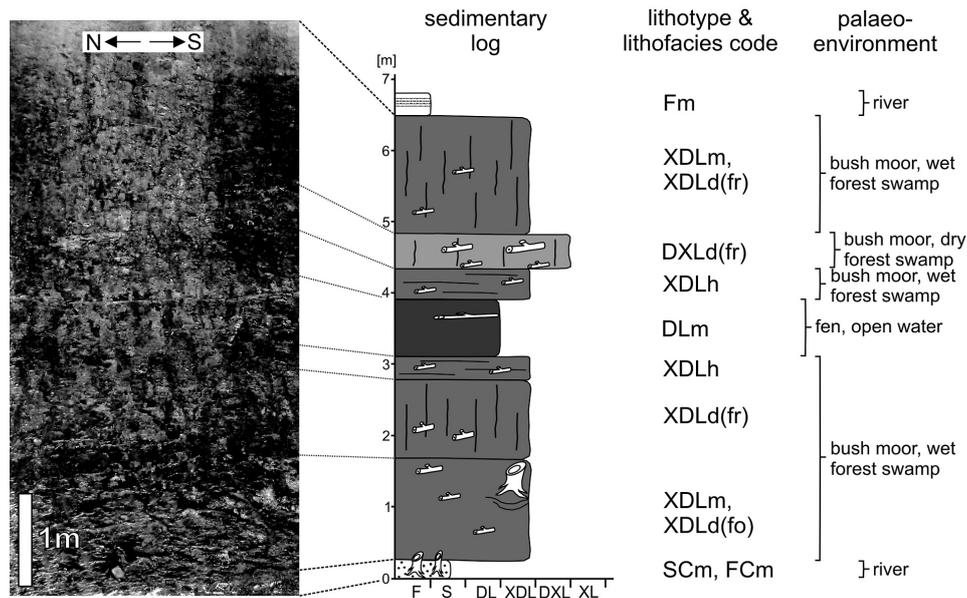


Fig. 6. Macroscopic vertical section of the 1<sup>st</sup> Middle-Polish Lignite Seam in the Koźmin N open-cast mine. Note the fold (lower-right corner of the section) formed by the weight of tree. For explanation of the lignite lithotype codes, see Table 3; for explanation of the lithofacies codes, see the text; for graphic explanations, see Figure 4.

Kwiecińska & Wagner (1997), Markič & Sachsenhofer (1997), Kolcon & Sachsenhofer (1999), Lücke et al. (1999), Ticleanu et al. (1999), Diessel et al. (2000), Markič et al. (2001), and others.

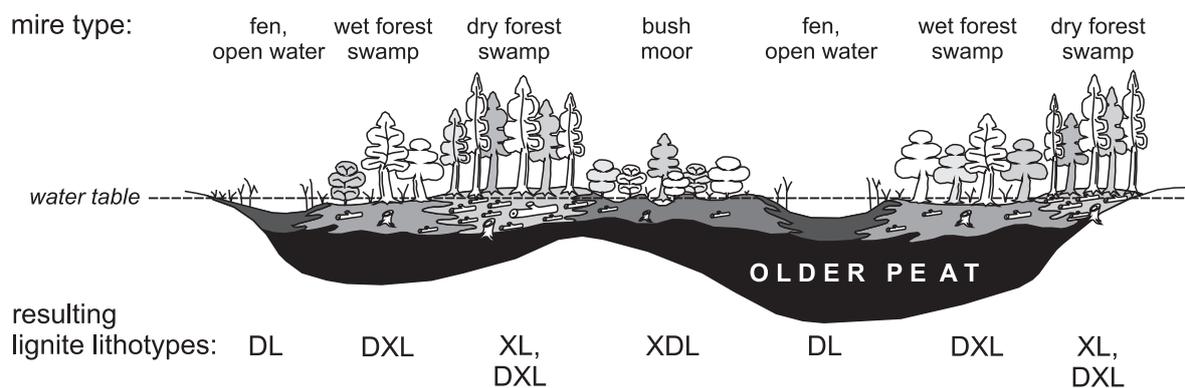
The first lithotype association is composed of one lignite lithotype, DLm (detritic lignite with massive structure). This lithotype is used as the prototype of a fen or open-water environment (Fig. 7). It may be related to a treeless reed marsh in the terminology of Teichmüller (1958, 1989). The sedges and reed-like vegetation as well as the aquatic plants should be regarded as the parent material for this lignite lithotype (Kolcon & Sachsenhofer, 1999; Ticleanu et al., 1999).

The second lithotype association, DXLh, DXLd(fr) and DXLd(fo), is characteristic of a wet forest swamp (Fig. 7). This type of a mire is also called the *Taxodium-Nyssa* swamp (Teichmüller, 1958, 1989). During the formation of this kind of mire the stable water table was relatively high above the mire surface (e.g. Lücke et al., 1999; Diessel et al., 2000). On the other hand, the xylites in growth position, mainly produced by a gymnosperm forest, point at a dry phase in the mire evolution (Kolcon & Sachsenhofer, 1999; Ticleanu et al., 1999).

The third lithotype association, XDLh and XDLd(fo), refers to a bush moor (e.g. Kolcon & Sachsenhofer, 1999; Fig. 7). It is also addressed as a *Myricaceae-Cyrrillaceae* swamp (Teichmüller, 1958, 1989). The peat-forming conditions were slightly drier than in the two above-mentioned associations (e.g. Lücke et al., 1999). During the formation of this lithotype association, the bushy plants occupied the wet areas (Kolcon & Sachsenhofer, 1999; Ticleanu et al., 1999).

The fourth lithotype association, XLm, XLd(fr) and XLd(fo), represents the dry forest swamp (Fig. 7). Taking into consideration the important role of *Sequoia*, this mire was named by Teichmüller (1958) a '*Sequoia* swamp'. These lithotypes, i.e. xylitic lignites with various structures, are the best record of dry conditions in the mire history. When they formed, the water table was relatively deep below the mire surface (e.g. Lücke et al., 1999; Diessel et al., 2000).

The occurrence of the above mire types in the MPLS-1 is partially supported by plant assemblages. The dry forest swamp is characterised, for example, by pollen of *Sequoia*, *Sciadopitys*, *Pinus*, *Betula* and *Fagus* (Sadowska & Giża, 1991; Ważyńska, 1998). In contrast,



**Fig. 7.** Mire types and resulting lignite lithotypes for the 1<sup>st</sup> Middle-Polish Lignite Seam in central Poland (modified after a model of Teichmüller, 1958, 1989). For explanation of the lignite lithotype codes, see Table 3; for graphic explanations, see Figure 4.

a predominance in the pollen spectrum of, for instance, *Taxodium*, *Nyssa* and *Alnus* indicates a wet forest swamp. Moreover, the depositional environment determined by *Cyrillaceae*, *Myricaceae*, *Ericaceae*, etc., is typical of a bush moor (Sadowska & Giża, 1991; Ważyńska, 1998).

It is noteworthy that the contact between the lignite seam and the overlying clastic deposits is always sharp (Figs 4–6). This may be a result of sudden changes in the depositional conditions, i.e. from a limnotelmatic or telmatic to a fluvial or fluviolacustrine environment (Piwocki & Ziemińska-Tworzydło, 1997; Widera, 2007). In contrast, the base of the lignite seam is sometimes gradual, e.g. in the Nazimierz N and Koźmin N sections. This may be explained by the presence of autochthonous stumps and small roots in the uppermost parts of the clastic deposits under the lignite seam (Figs 4, 6).

The MPLS-1 in central Poland developed while a land (fluvial environment) was replaced by a mire (limnotelmatic or telmatic environments). This transition process is called 'paludification' (Diessel et al., 2000). However, the lignite sections differ in detail. The lignite seam from the Nazimierz N mine, for example, shows a drying- and shallowing-upward trend. It is characterised by the lithotype transition from massive detritic lignite (DLm) to massive xylitic lignite (XLm) (Fig. 4). In contrast, a clear upward trend, showing either drying or wetting, is not traceable in the other lignite sections. This is shown, for example, in the Drzewce and Koźmin N sections, where the driest

xylitic lithotypes with various structures, i.e. XLd(fo) and XLm, are present in the middle part of the lignite seams (Figs 5–6).

Deformational structures such as folds and fractures are also common in the MPLS-1 in central Poland. The folds within the lignite seams may be interpreted in at least three ways. They can have been formed by tectonics (e.g. Ticleanu et al., 1999), by compaction (e.g. Widera et al., 2007) and due to the weight of trees. In the last-mentioned case, the folds were caused by the weight of peat-forming trees (Figs 3–5). The fractures in the lignite seams are interesting and important structures for the mining activity (Figs 2, 4, 6). Their origin is still under discussion, but commonly they are attributed to a combination of tectonic and diagenetic processes (Laubach et al., 1998).

On the basis of all above-mentioned lignite properties (lithotypes, thickness, ash content, tectonic background, etc.) a depositional model of the MPLS-1 can be discussed. The development of the mires occurred in tectonics-induced subsidence (Widera, 2007; Widera et al., 2007). On the other hand, the lithotype associations, the plant assemblages and the ash content (> 10%) point at low-lying mires according to McCabe (1984) or to topogenous, rheotrophic mires according to Diessel et al. (2000). Such mire areas probably were periodically covered by shallow (< 2 m) water (Diessel et al., 2000). This is supported by the presence of the detritic lithotypes in the various sections (Figs 4–6).

## 7. Conclusions

The lithotype composition of the 1<sup>st</sup> Middle-Polish (1<sup>st</sup> Lusatian) Lignite Seam (MPLS-1) was investigated on the basis of three vertical sections in open-cast mines in central Poland. Each of these mines is located in a graben, which subsided during the development of the mires in the Middle Miocene.

Each lithotype shows specific physical and chemical properties which reflect the depositional environment.

Four lignite lithotype associations have been identified in the MPLS-1. They are the detritic, the xylo-detritic, the detro-xylic and the xylic associations, each with various individual structures. Each of them may be related to at least one mire type: fen or open water, bush moor, wet forest swamp and dry forest swamp, respectively (Fig. 7). The sections under study show that the mires formed as a result of paludification.

The investigation of lignite seams requires practical and unambiguous descriptions of the lithotypes in the field. A new and more tight codification of lignite lithotypes is therefore proposed here. The proposed codes for lignite lithotypes can be applied together with those for lithofacies, in the same way as currently commonly done for sedimentological descriptions.

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